

Spatio-temporal weather generator for the temperature over France

Présentation au séminaire de statistiques de Rochebrune

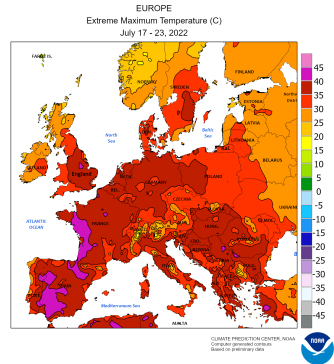
Caroline Cognot (MIA-PS & EDF)

Supervisors : Liliane Bel (MIA-PS) et Sylvie Parey (EDF)

Electricity and climate studies

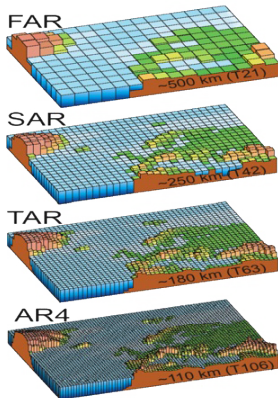


- Climate extremes, risks : impact on agriculture, health, energy production and demand
- Climate change : impacts frequency and intensity of spatial = multivariate meteorological hazards
- In particular, both on electricity generation and system balance



Studies are necessary and we use models to do this.

Physical climate modelling



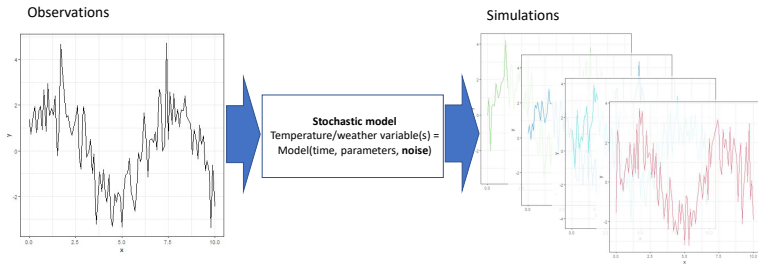
Climate models grid sizes

- Used in RCP modelling (regional models)
- Complex phenomena, well-known from physics

BUT

- Computationally expensive
- Not adapted to extremes

Why stochastic modelling



- Simulations = same statistical properties as the observations
- Reproduce the properties important to the user (e.g. heat waves)
- Computationally efficient
- Can be used to sample the distribution, including extremes
- Can be used with real data or debiased climate model output

"Stochastic Weather Generator"

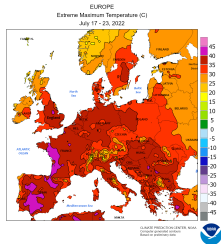
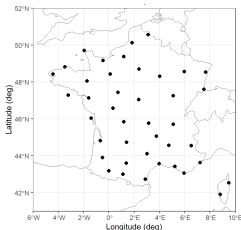
A temperature generator

Existing work :

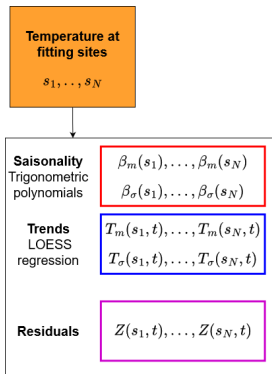
- Single-site models : models for $n \geq 1$ variables $X^1(t), \dots, X^n(t)$
- Multi-site models, but viewing sites as many different variables : models for $n \geq 1$ variables at $n_S \geq 1$ sites $X_1^1(t), \dots, X_1^n(t), \dots, X_{n_S}^1(t), \dots, X_{n_S}^n(t)$
- Spatial models, focused mostly on precipitations : models for $n \geq 1$ variables on a spatial extent $\{s \in \mathcal{D}\}$ $X^1(s, t), \dots, X^n(s, t)$

My objective :

Build a SWG for the temperature, reproducing the spatial and temporal structure and allowing for sampling in the climate variability.



Decomposition in trend and seasonality¹



Objective : separate deterministic from stochastic components. Deterministic = climate, stochastic = climate variability.

For each site s and time t , we decompose the temperature $X(s, t)$:

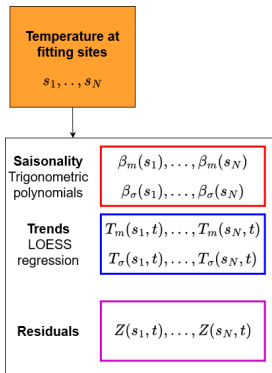
$$X(s, t) = T_m(s, t) + S_m(s, t) + T_\sigma(s, t) S_\sigma(s, t) Z(s, t)$$

where

- $T_m(s, \cdot)$ is the long-term **mean trend**;
- $T_\sigma(s, \cdot)$ is the long-term **variance trend** ;

¹Hoang 2010.

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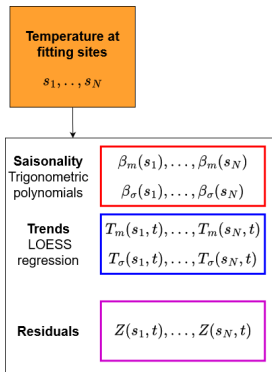
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- $T_m(s, \cdot)$ is the long-term **mean trend**;
- $T_\sigma(s, \cdot)$ is the long-term **variance trend** ;
- $S_m(s, \cdot)$ is the **mean seasonality**;
- $S_\sigma(s, \cdot)$ is the **variance seasonality**;

$$S_{m \text{ or } \sigma}(s, t) = \beta_1(s) + \sum_{i=1}^{d_m \text{ or } \sigma} \left[\beta_{2i}(s) \cos\left(\frac{2i\pi t}{365}\right) + \beta_{2i+1}(s) \sin\left(\frac{2i\pi t}{365}\right) \right]$$

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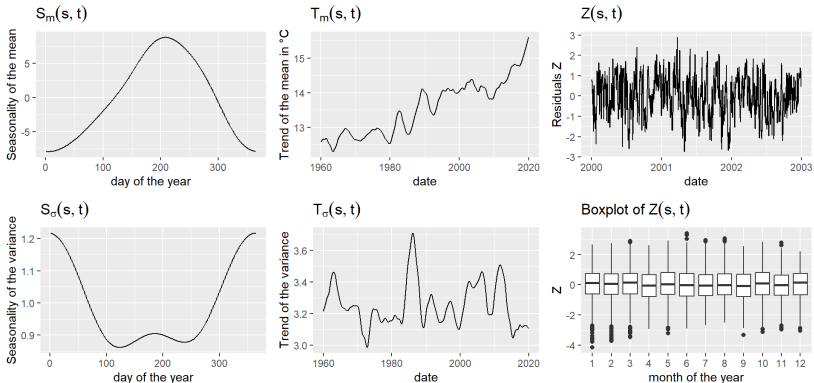
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- $Z(s, \cdot)$ are the **residuals** at site s . They have variance 1 and mean 0.

¹Hoang 2010.

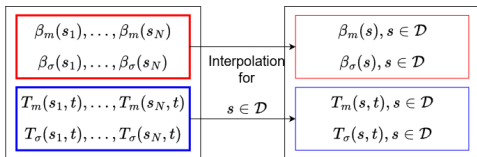
Example : s fixed, Toulouse

$$X(s, t) = T_m(s, t) + S_m(s, t) + T_\sigma(s, t)S_\sigma(s, t)Z(s, t)$$

Decomposition of the temperature - Toulouse-Blagnac

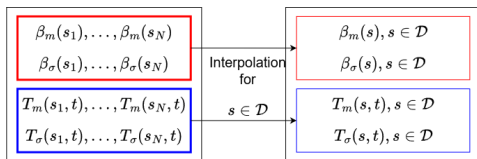


Decomposition : spatial extension



Objective :
Extend decomposition
parameters from discrete to
continuous
 \mathcal{D} is a high resolution grid

Decomposition : spatial extension

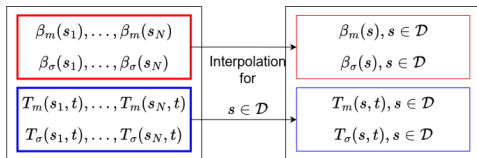


Objective :
Extend decomposition parameters from discrete to continuous
 \mathcal{D} is a high resolution grid

- For trends : Trends are "smooth" in time and slow-varying : reasonable to use constant weights, using Inverse Distance Weighting (IDW) interpolation \rightarrow gives many maps of slow-varying mean temperature and standard deviation.

$$u(s) = \frac{\sum_{i=1}^n w_i(s) u_i}{\sum_{i=1}^n w_i(s)}, \text{ with } w_i(s) = \frac{1}{(d(s_i, s))^p}$$

Decomposition : spatial extension



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- For **trends** : Trends are "smooth" in time and slow-varying : reasonable to use constant weights, using Inverse Distance Weighting (IDW) interpolation \rightarrow gives many maps of slow-varying mean temperature and standard deviation.
- For **seasonality** : kriging on the coefficients \rightarrow 2 maps of spatial coefficients, to multiply by corresponding sines and cosines to obtain the cycle.

Model of the residuals

Recall : decomposition

$$X(s, t) = T_m(s, t) + S_m(s, t) + T_\sigma(s, t)S_\sigma(s, t)Z(s, t)$$

We want to model the stochastic part $Z(s, t)$, which is already supposed to be zero-mean, with variance close to 1. :

Proposal : 2nd-order spatio-temporal model

A 2nd-order spatiotemporal model $Z(\mathbf{s}, t)$ is defined by its 2 components :

1. The mean function $\mathbb{E}[Z(\mathbf{s}, t)] = \mu(\mathbf{s}, t) = 0$ for us
2. The covariance function $C(\mathbf{s}_1, \mathbf{s}_2, t_1, t_2) = \text{Cov}[Z(\mathbf{s}_1, t_1), Z(\mathbf{s}_2, t_2)]$

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Model of the residuals

We want to model the stochastic part $Z(\mathbf{s}, t)$, which is already supposed to be zero-mean, with variance close to 1. :

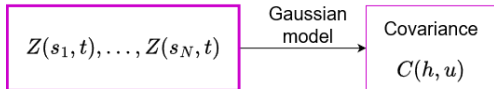
Proposal : 2nd-order stationary **isotropic** spatio-temporal model

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This reduces our problem to fitting a covariance function

$$C: \mathbb{R}^+ \times \mathbb{R} \longrightarrow \mathbb{R}$$
$$(h, u) \longmapsto C(h, u)$$



Model of the residuals : Possible covariance functions

- Separable space-time functions : $C(h, u) = C_S(h)C_T(u)$:

²Gneiting 2002.

Model of the residuals : Possible covariance functions

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Model of the residuals : Possible covariance functions

- Separable space-time functions : $C(h, u) = C_S(h)C_T(u)$: no space-time interactions !
- Non - separable functions : many classes of Gneiting type².
For "nice" $\phi(t)$ and $\psi(t)$

$$C(h, u) = \frac{\sigma^2}{\psi(|u|^2)^{d/2}} \phi\left(\frac{\|h\|^2}{\psi(|u|^2)}\right)$$

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Matérn kernel : $C(h, u) = \frac{\sigma^2}{(\alpha u^{2a} + 1)^b} \mathcal{M}\left(\frac{h}{\sqrt{\alpha u^{2a} + 1}}; r; \nu\right)$

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In practice, multiplied by a purely temporal covariance function for the Gneiting-Matérn class :

Gneiting-Matérn covariance function

$$C(h, u) = \frac{\sigma^2}{(\alpha u^{2a} + 1)^{b+\delta}} \mathcal{M}\left(\frac{h}{\sqrt{\alpha u^{2a} + 1}}; r; \nu\right)$$

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Model of the residuals : Parameter estimation

Gneiting-Matérn covariance function

$$C(h, u) = \frac{\sigma^2}{(\alpha u^{2a} + 1)^{b+\delta}} \mathcal{M} \left(\frac{h}{\sqrt{\alpha u^{2a} + 1}^b}; r; \nu \right) \rightarrow 7 \text{ parameters.}$$

log-likelihood : for $(X_1, \dots, X_n) \sim \mathcal{N}(0, \Sigma)$

$$\begin{aligned} \ell(x, \theta) &= \log(f(X_1 = x_1, \dots, X_n = x_n)) \\ &= -\frac{1}{2} \log(|\Sigma|) - \frac{1}{2} ((x_1 \dots x_n) \Sigma^{-1} (x_1 \dots x_n)^T) \end{aligned}$$

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Here, $n = 41 \times 365$ for a year of data ! Inversion is not recommended.

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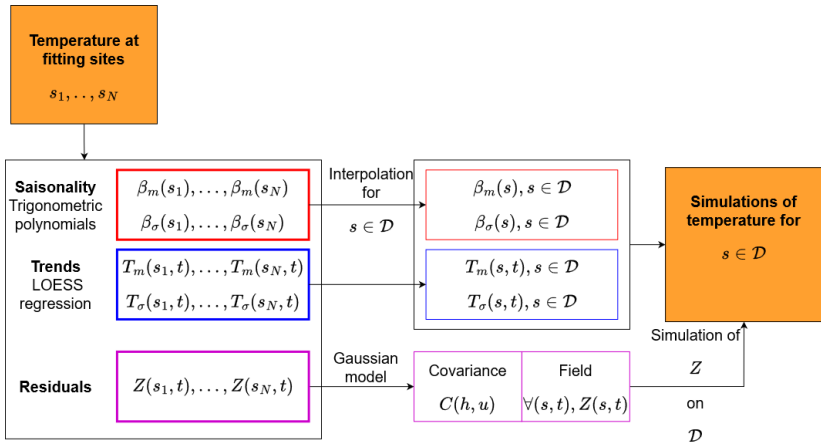
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Outline of the model



Simulation methods for the Gaussian field

- Naive approach : simulate from multivariate Gaussian covariance matrix

³Allard et al. 2020.

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Not feasible for large space and time grid : **ok for 40 points in space, but not for 1000.**



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Not feasible for large space and time grid : **ok for 40 points in space, but not for 1000.**
- A bit more refined : simulate all spatial points for each day t from previous l days using Gaussian properties and stationnarity : for X_i the vector at all points in space at time i , there are matrices A, B, C such that

$$\begin{bmatrix} X_1 \\ X_2 \\ \dots \\ X_l \end{bmatrix} \sim \mathcal{N}(0, A)$$

$$\forall t > l, X_t \mid \begin{bmatrix} X_1 \\ \dots \\ X_{t-1} \end{bmatrix} = X_t \mid \begin{bmatrix} X_{t-l} \\ \dots \\ X_{t-1} \end{bmatrix} \sim \mathcal{N} \left(B \begin{bmatrix} X_{t-l} \\ \dots \\ X_{t-1} \end{bmatrix}, C \right)$$

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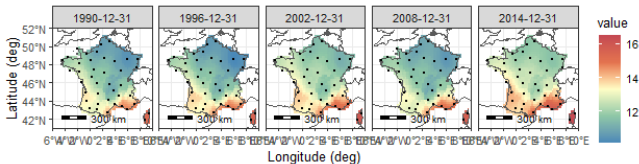
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- A specific algorithm for Gneiting-type covariance functions : use the spectral algorithm³

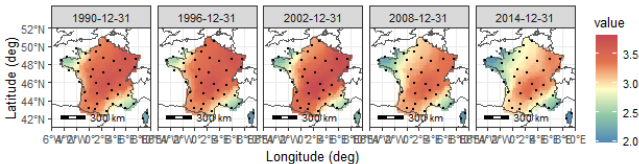
³Allard et al. 2020.

Decomposition : trends

Trend in mean - IDW interpolation



Trend in variance - IDW interpolation

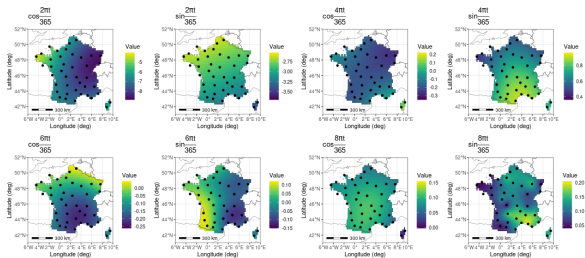


- Choice of presented results : same date each year, but there is no monotonicity (see Toulouse result earlier, especially for the variance)
- Trends show expected French climate

Decomposition : seasonality of the mean

Recall : Seasonality

$$S_{m \text{ or } \sigma}(s, t) = \beta_1(s) + \sum_{i=1}^{d_m \text{ or } \sigma} [\beta_{2i}(s) \cos\left(\frac{2i\pi t}{365}\right) + \beta_{2i+1}(s) \sin\left(\frac{2i\pi t}{365}\right)]$$

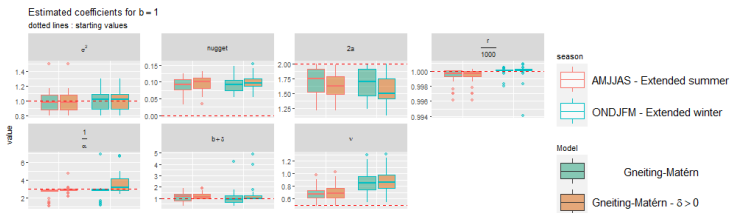


- Different "frequency" give different maps
- Points out different climates
- What could be useful : instead of $A \cos(2k\pi t/365) + B \sin(2k\pi t/365)$, have $C \cos(2k\pi t/365 + \Phi)$ can give more temporal insight
- Seasonality in variance is less interesting to show

Model of the residuals : estimated parameters

Recall : Gneiting-Matérn covariance function

$$C(h, u) = \frac{\sigma^2}{(\alpha u^{2a} + 1)^{b+\delta}} \mathcal{M} \left(\frac{h}{\sqrt{\alpha u^{2a} + 1}}^b; r; \nu \right)$$

Grid search for b in $[0, 1]$ → maximum of likelihood for 1.

- Spatial range parameter stayed at initial value : probable redundancy with α, a, b inside the Matérn kernel
- Winter and summer have different ν (smoothness parameter)

Validation

Idea : Compute indicators from the observations and compare with the same indicators from many simulations. The model is adequate if the observed values are in the range of the simulations.

Indicators of good fit :

- **Pairwise correlation** between pairs of stations
- **Pairwise conditional threshold exceedance**
For every pair of stations i, j , define

$$p_{i,j}^{\alpha} = P(X_i > q_{\alpha}(i) | X_j > q_{\alpha}(j))$$
$$\hat{p}_{i,j}^{\alpha} = \frac{\sum_{t=1}^{N_t} \mathbf{1}_{X_i > q_{\alpha}(i) \cap X_j > q_{\alpha}(j)}}{\sum_{t=1}^{N_t} \mathbf{1}_{X_j > q_{\alpha}(j)}}.$$

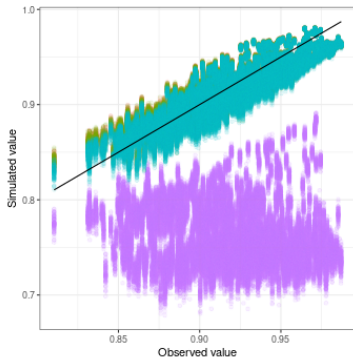
With inverted signs in the case of low quantiles.

- **Lagged temporal auto-correlation**

Validation : simulation at the fitting points

Pairwise correlation between stations

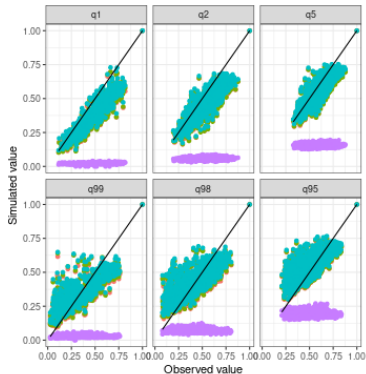
all simulated values for each pair



$P(Y_i > q_i^{obs} | Y_j > q_j^{obs})$ for high quantiles

$P(Y_i < q_i^{obs} | Y_j < q_j^{obs})$ for low quantiles

Median simulated values for each pair

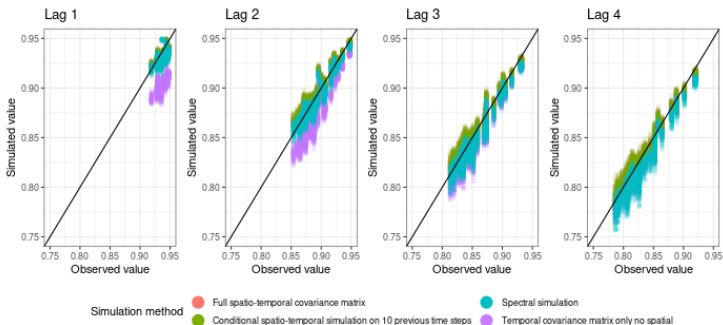


Simulation method

- Full spatio-temporal covariance matrix
- Conditional spatio-temporal simulation on 10 previous time steps
- Spectral simulation
- Temporal covariance matrix only no spatial

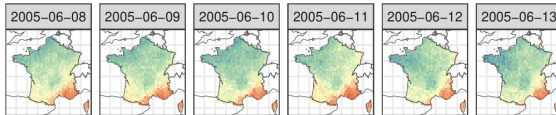
Validation : simulation at the fitting points

Temporal correlation

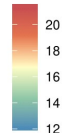


Simulation on a grid : what it looks like

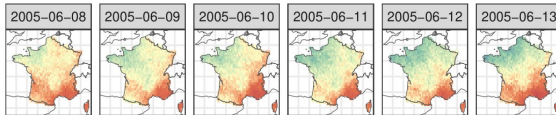
Simulation 1



Temperature (°C)



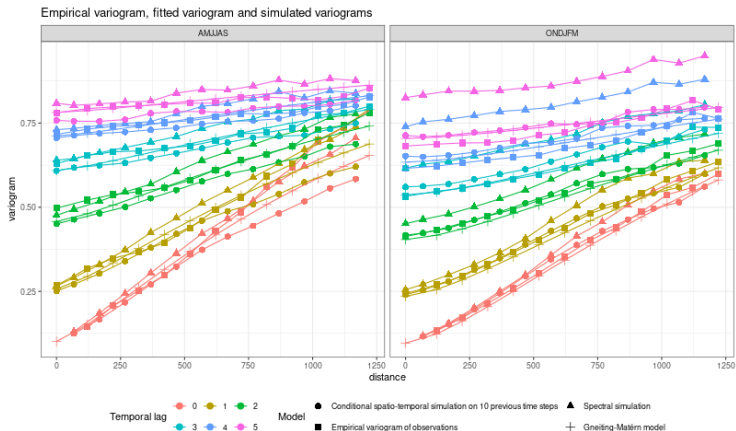
Simulation 2



180 days of simulation = 150s (spectral method, can be improved) or 10s (conditionnal iterative method, but with additional time for matrix creation)

Simulation on a grid : variograms

$$\text{Variograms : } \gamma(h, u) = \frac{1}{2} \text{Var}(Z(s + h, t + u) - Z(s, t))$$



- The theoretical model is close to the observations
- The simulations are close to the model (the methods worked)

Conclusion

What is done :

- A model for temperature that takes into account spatial structure
- Results : spatial correlation well reproduced, extremal dependence not so much (but not so bad). May need refining.

What I want to do next :

- Compare simulations on a grid to EOBS dataset (same grid)
- Compare interpolated decomposition with exact one
- Use this model with a precipitation model